

Importance of marine boundary layer clouds for the mean climate and interannual variability over the Atlantic Ocean

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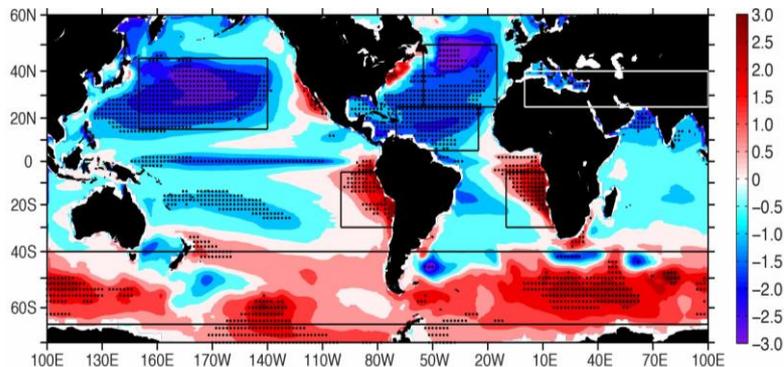
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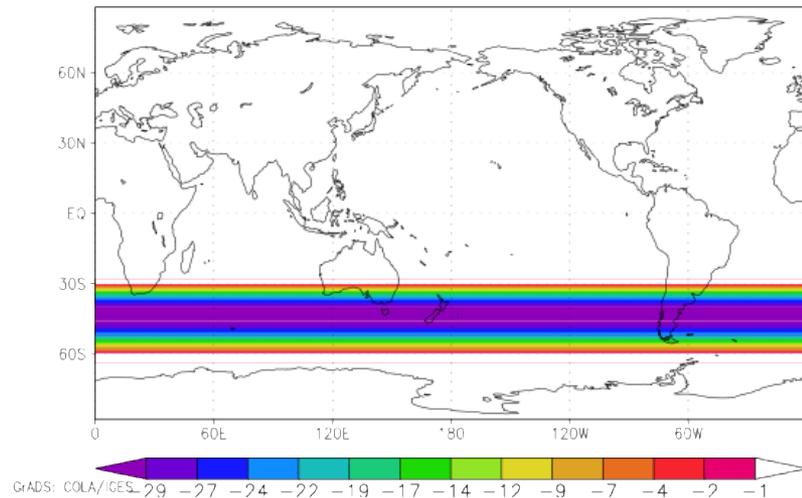
⁴ U. de la Republica, Uruguay

SST biases averaged in CMIP5 models



Wang, Zhang, Lee, Wu, and Mechoso
(2014, *Nature CC*)

Reduction in Incident SW at TOA



Questions and Approaches

Question 1: Can reducing the incoming energy flux over the southern ocean in a CGCM improve its simulation of tropical climate?

Approach: Contrast the effects of reducing SW incident at TOA over the Southern Ocean in two different models: **UCLA CGCM** and **NorESM**

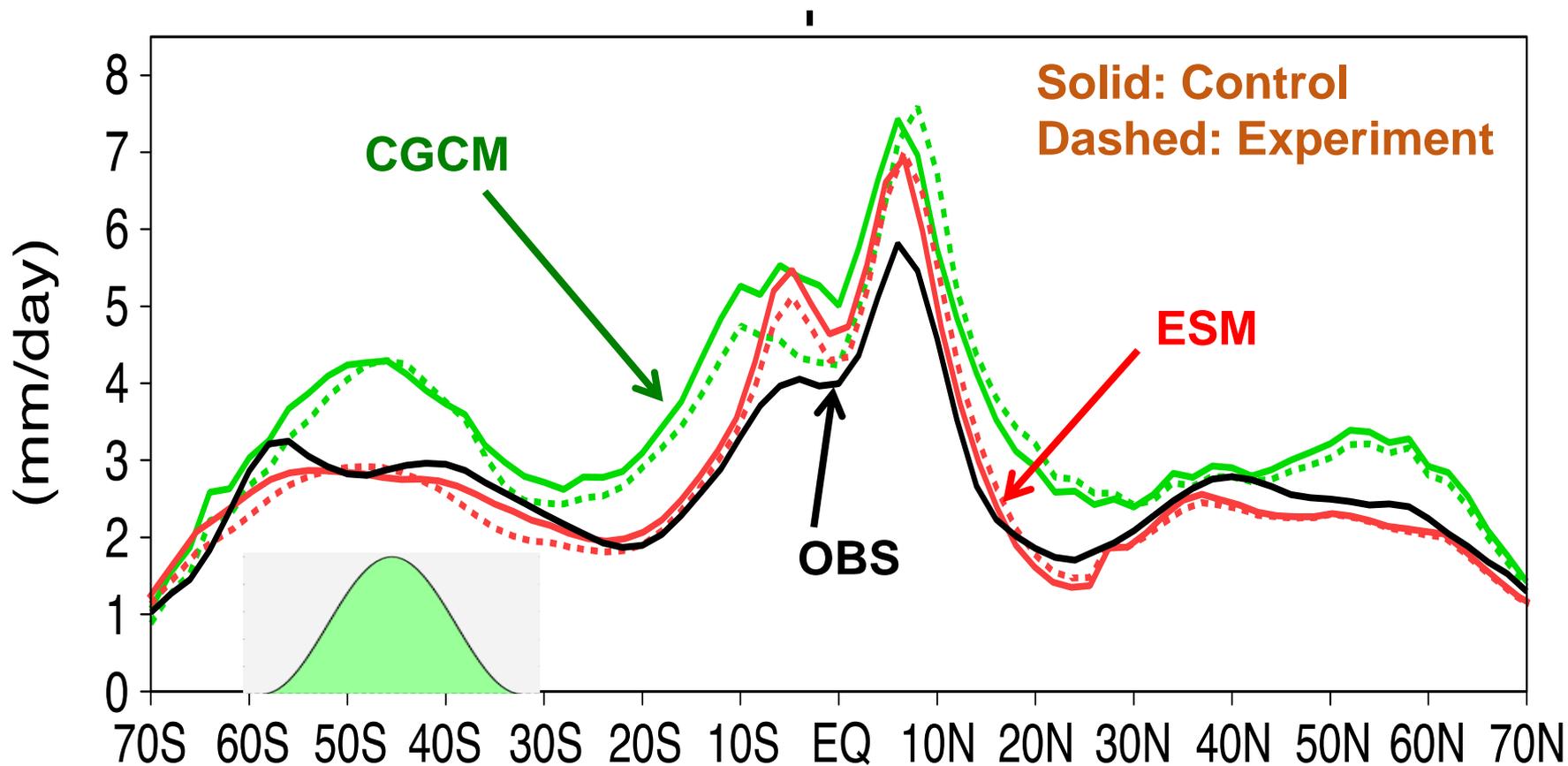
(Hereafter **CGCM** and **ESM**)

Question 2: Is SST variability over Atlantic Ocean amplified by positive cloud feedbacks?

Approach: Examine the regression of detrended seasonal anomalies in SW cloud radiative effect and in SST onto mean SST anomalies in the Scu regions of the eastern Atlantic.

(Use **ISCCP**, **CERES**, **ERA-Interim** and **CMIP5** data)

Annual and zonal mean precipitation show less double ITCZ

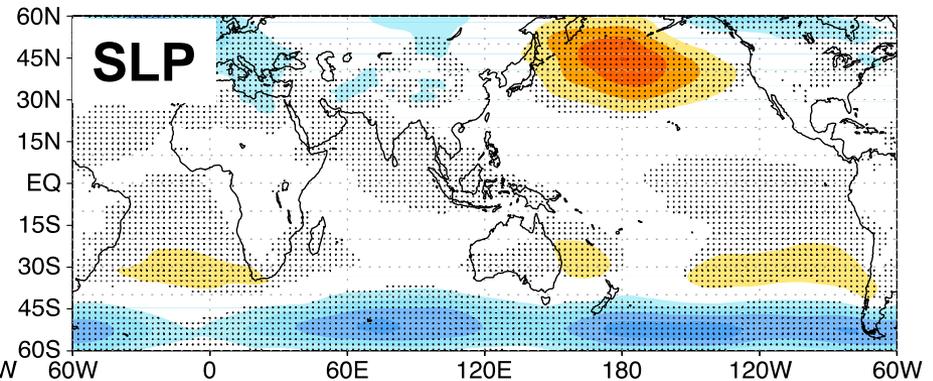
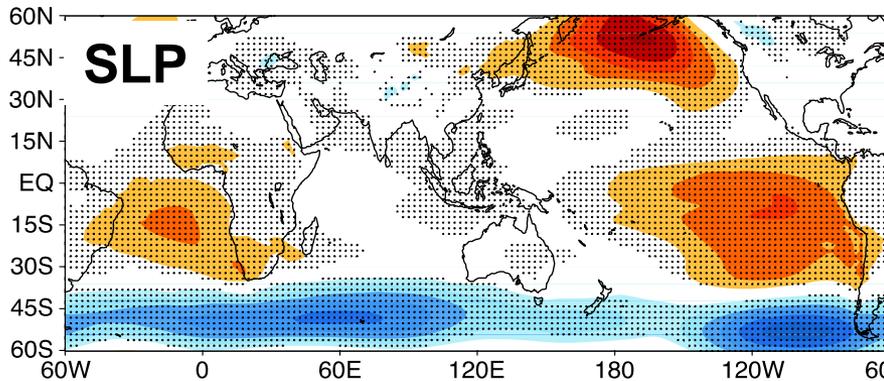
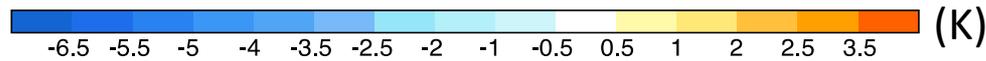
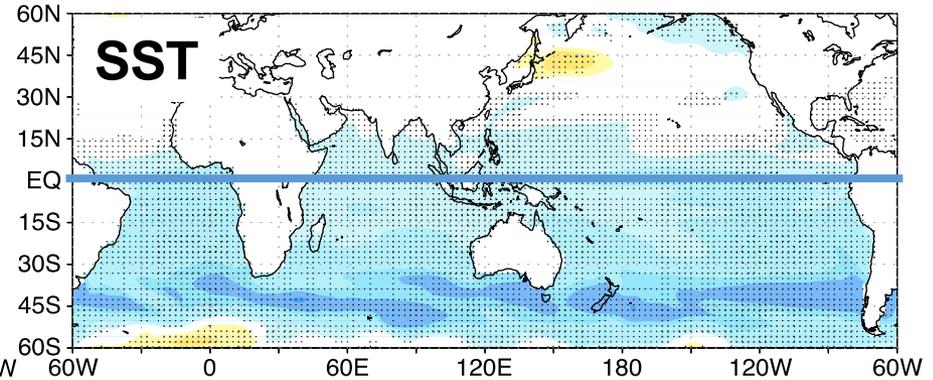
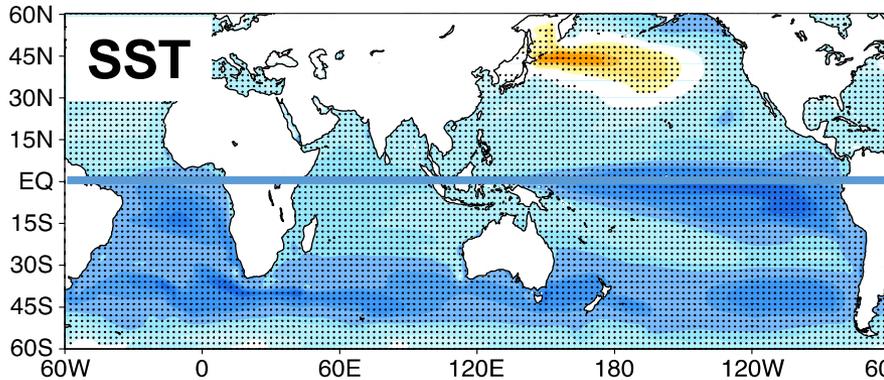


Short-Wave Incident at TOA
is decreased between 30S-
60S ~ 3.7% rad received
between 20S and 90S

SSTs cool down and the southern subtropical highs are enhanced

CGCM

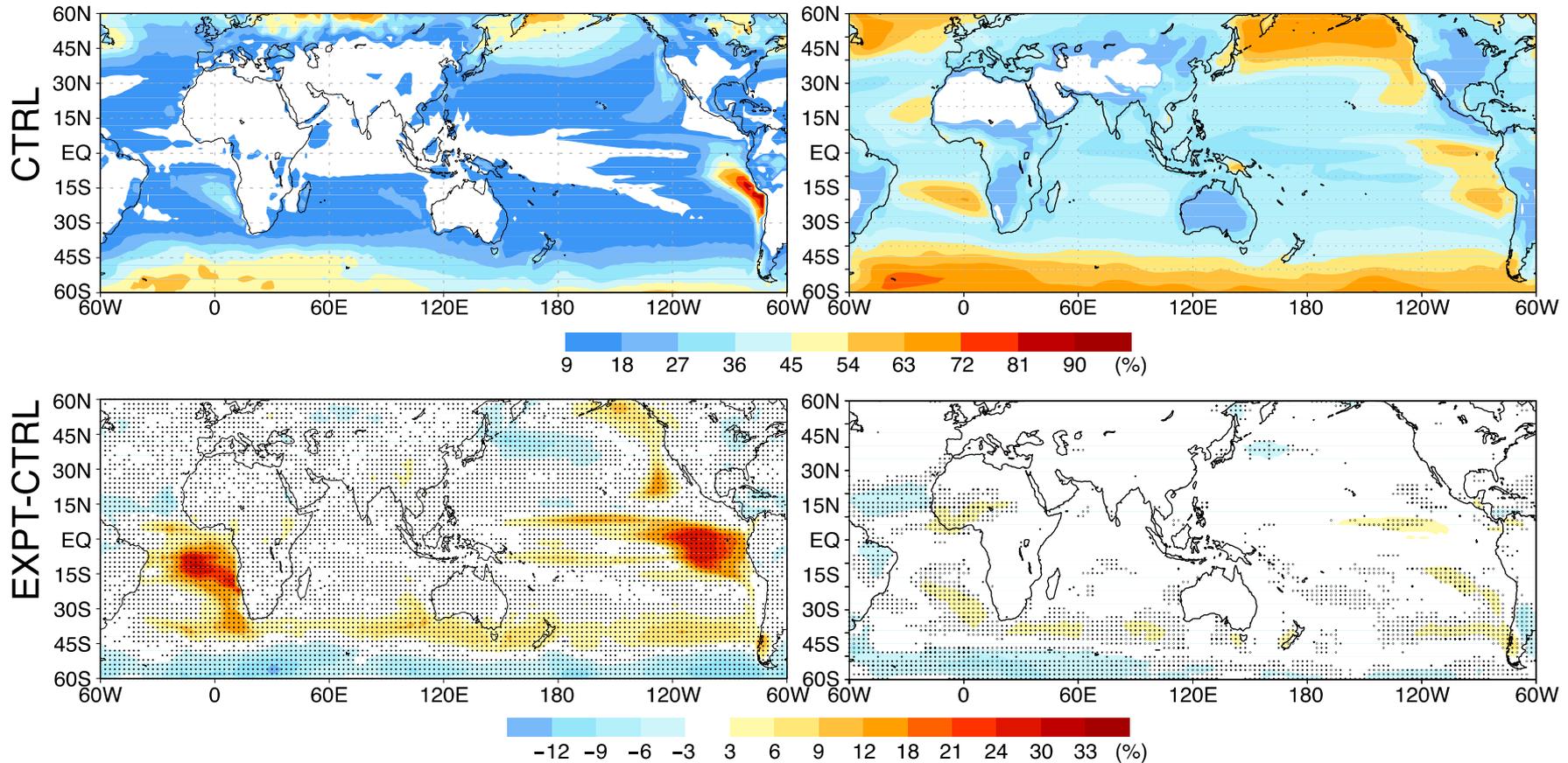
ESM



Changes in TOA radiation are consistent with changes in low-level clouds

CGCM Scu (PBL) clouds

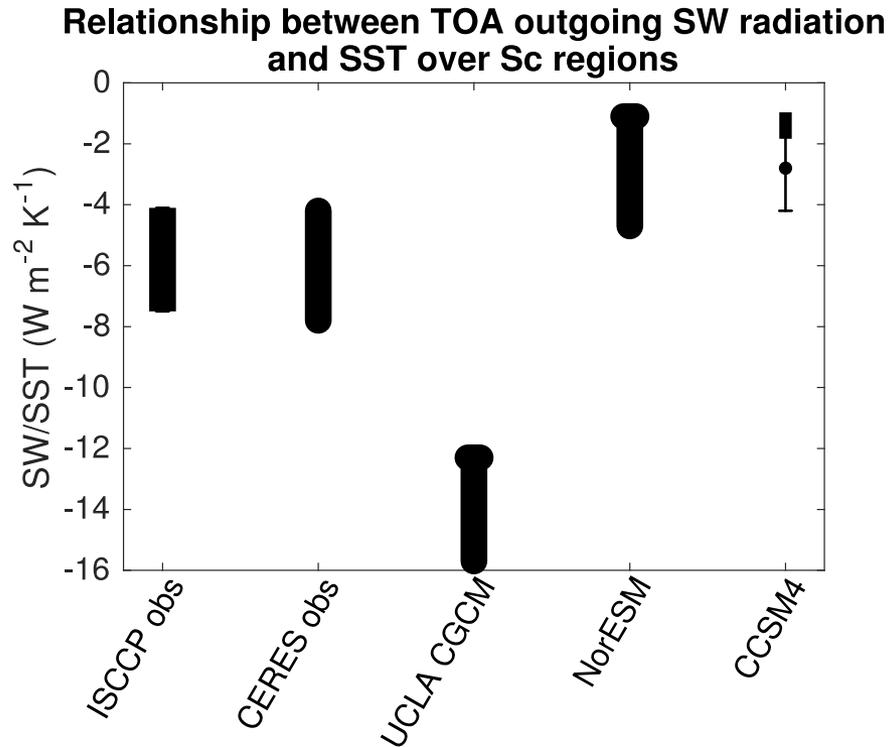
ESM Low-level clouds



Top: Control

Bottom: Experiment-Control

Slope of regression of monthly anomalies of outgoing SW radiation on SST over the five major stratocumulus regions over the global oceans.



Errorbars denote 95% confidence bounds, taking into account temporal and spatial autocorrelation

Data

Cloud radiative effect (CRE) = clear-sky minus all-sky outgoing radiation at top of atmosphere

Cloud fraction

ISCCP 1985-2000

CERES 2001-2014

SST

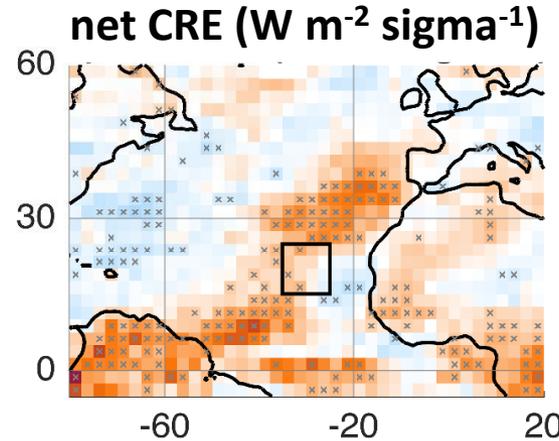
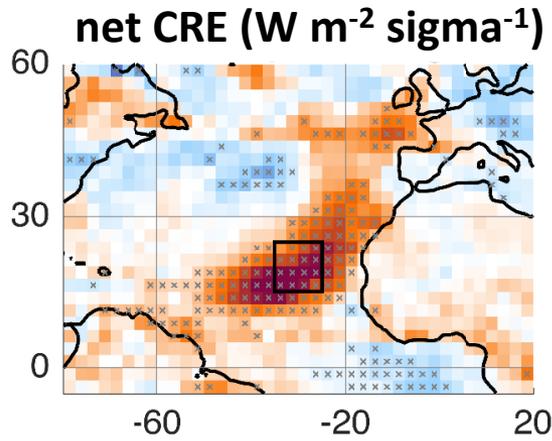
Sea-level pressure (SLP), winds

ERA-Interim reanalysis

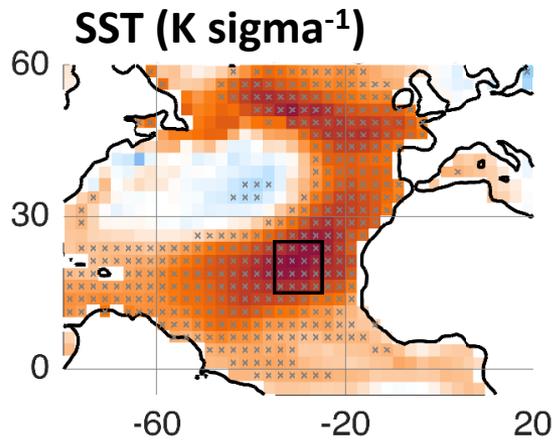
(same results using NOAA Optimum Interpolation SST V2)

CMIP5 historical runs 1976-2005

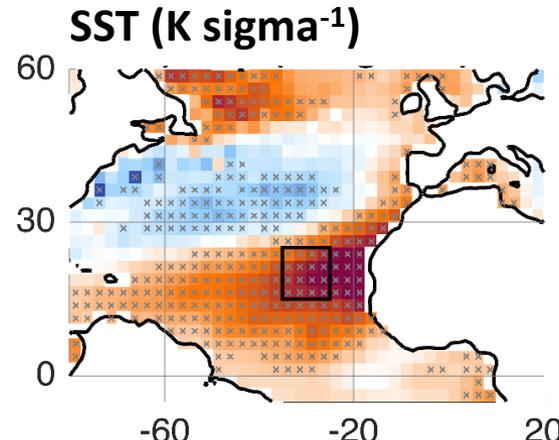
Patterns associated with dominant mode of North Atlantic SST variability



Summer JJA:
Consistent
w/positive
cloud feedback



JJA

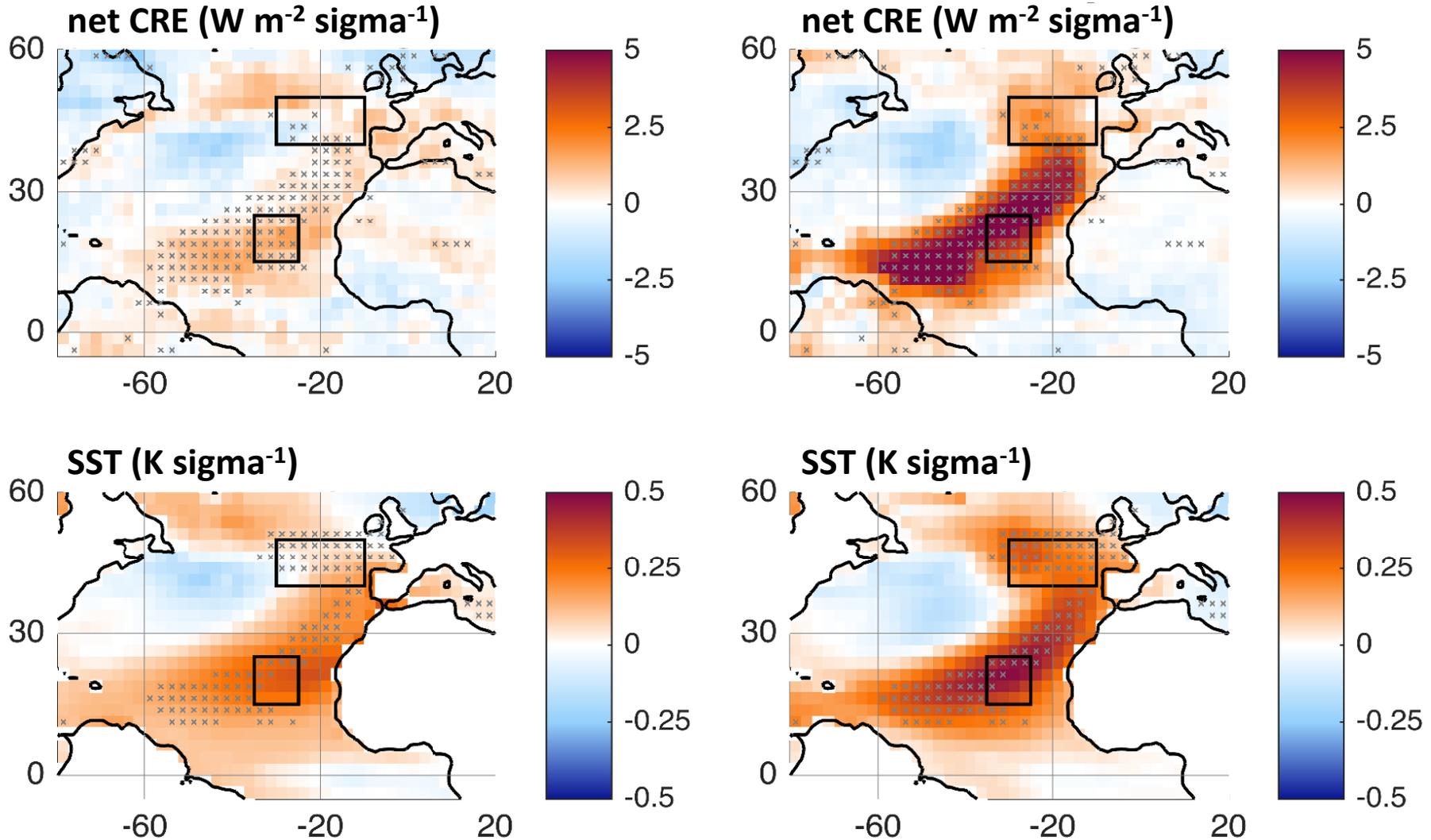


DJF

Winter DJF:
Inconsistent
w/positive
cloud feedback

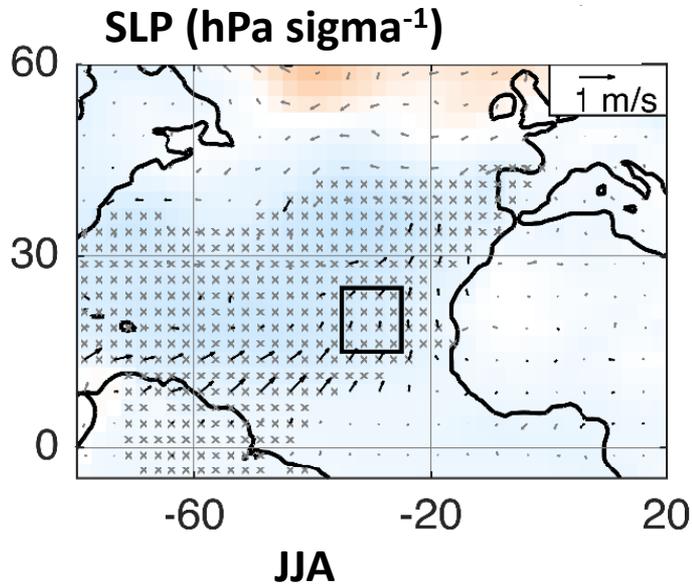
Slopes of regression of **net cloud radiative effect (CRE)** and **SST** seasonal anomalies onto 1985-2014 SST anomalies averaged over boxed region of max boundary layer cloud amount.

Patterns of dominant mode of summertime North Atlantic SST variability in models with weak/strong cloud feedback over subtropical NE Atl.



Statistically significant basin-scale difference in net CRE and SST

North Atlantic Oscillation (NAO) emerges in winter but not in summer

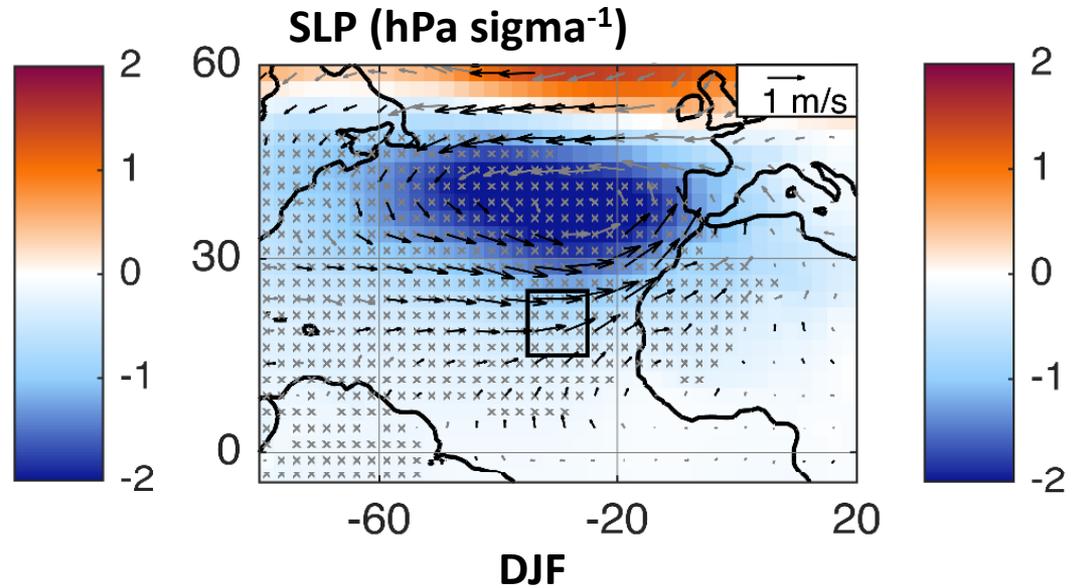


i) weaker heat flux anomalies

ii) shallower ocean mixed layer

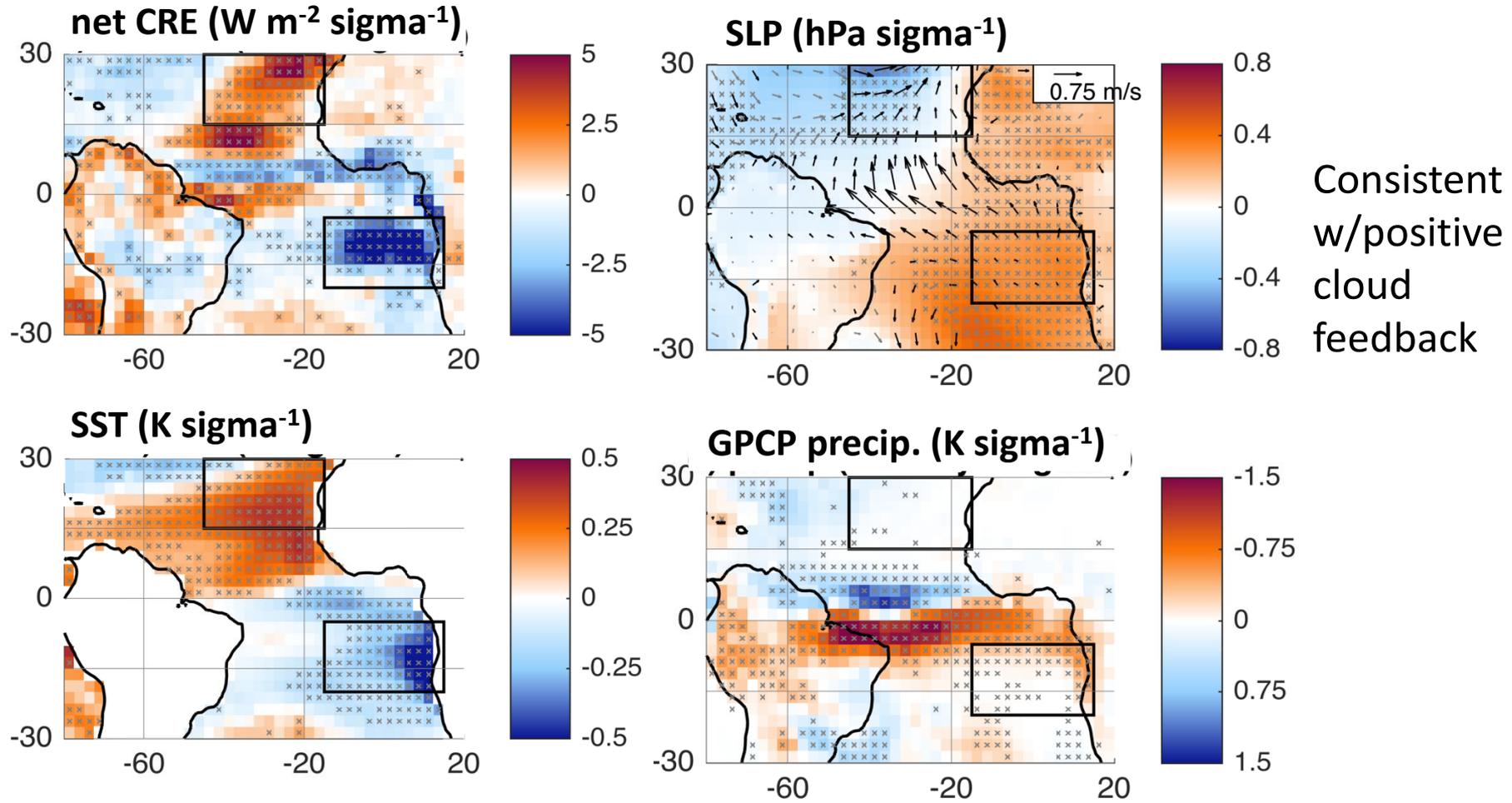
greater role for clouds

Slopes of regression of **sea-level pressure (SLP)** and **wind** seasonal anomalies onto 1985-2014 SST anomalies averaged over boxed region of max boundary layer cloud amount.



Surface winds drive sensible and latent heat fluxes that generate SST anomalies.

Patterns associated with dominant mode of tropical Atlantic coupled variability during spring



Slopes of regression of climate field MAM anomalies onto 1985-2014 SST anomaly difference between boxes.

Conclusions

Question 1: Can reducing the incoming energy flux over the southern ocean in a CGCM improve its simulation of tropical climate?

Yes. The extent of improvement depends upon the CGCM's success in capturing Scu-SST feedbacks.

Question 2: Is SST variability over Atlantic Ocean amplified by positive cloud feedbacks?

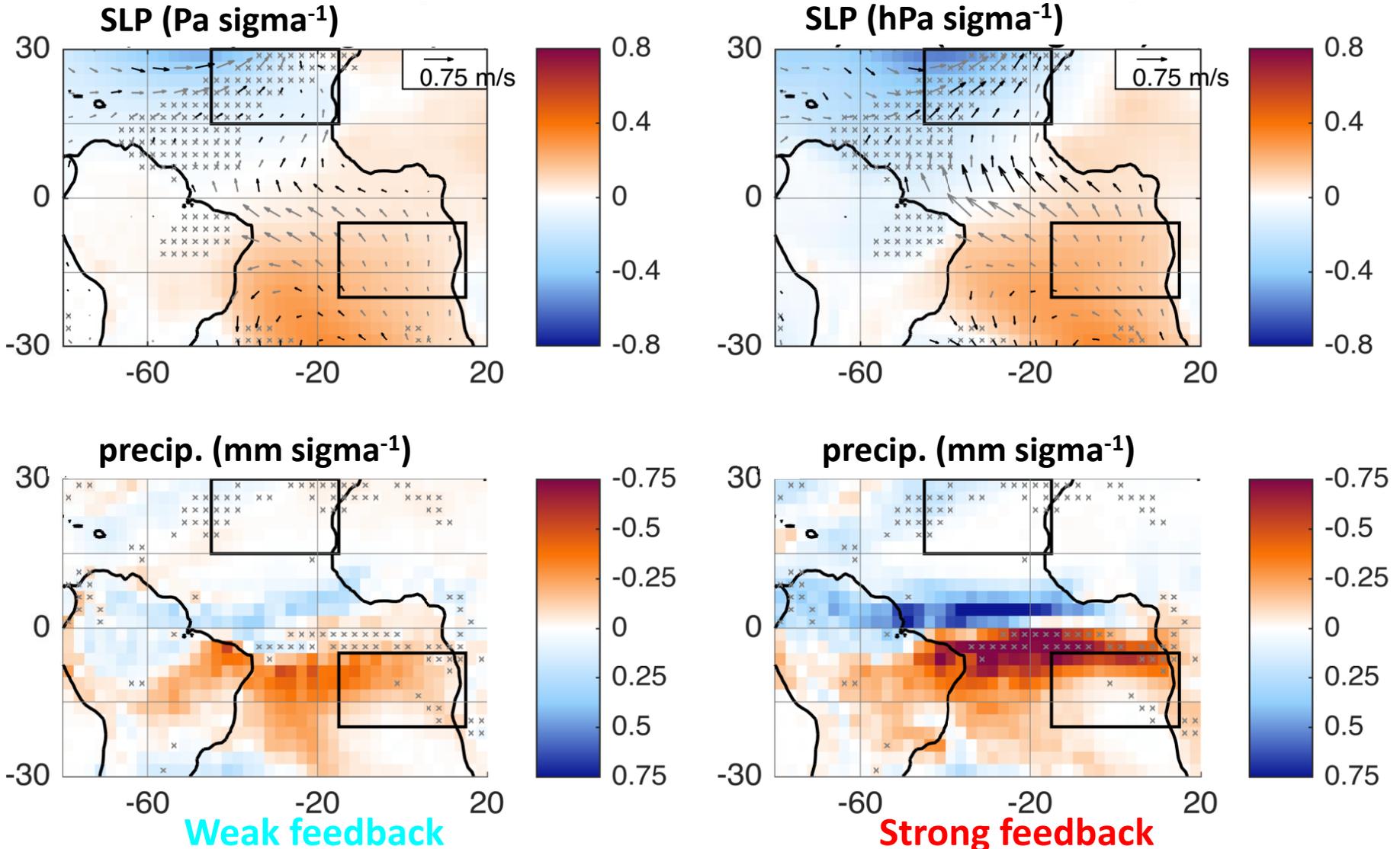
Yes. On the basis of observational data, we find new evidence for linkage between cloud-SST feedback and dominant modes of variability.

Conclusions

Improve marine boundary layer clouds

**→ more realistic simulation of mean
climate and interannual to
interdecadal atmosphere-ocean
variability**

Patterns of dominant mode of springtime tropical Atlantic coupled variability in CMIP5 models with weak/strong cloud feedback over subtropical SE Atlantic



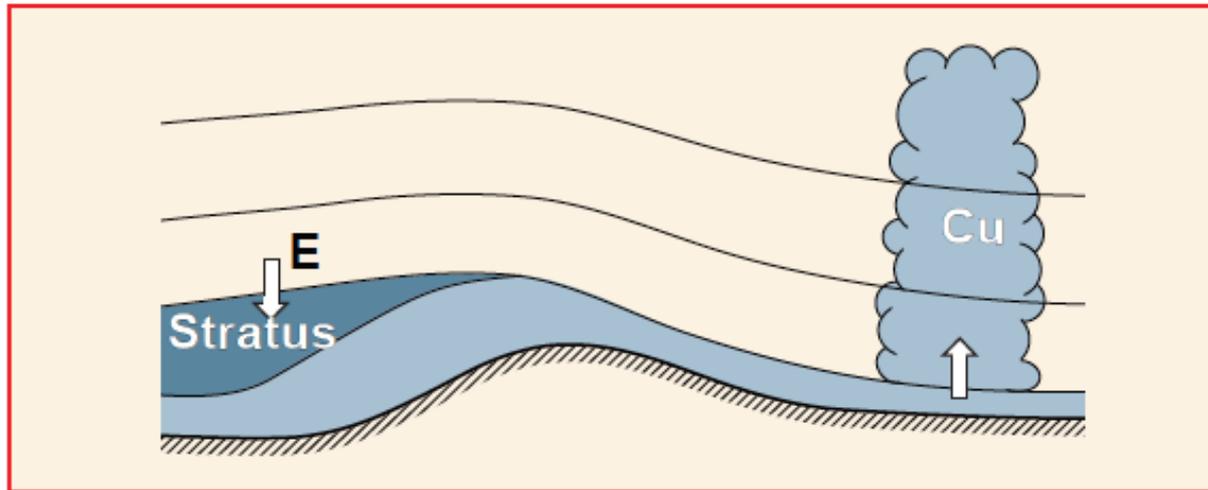
Weak feedback

Strong feedback

Statistically significant difference in SLP, winds, and precipitation

PBL PARAMETERIZATION IN UCLA AGCM

Suarez, Randall and Arakawa(1983) (Gen IV-V)



Characteristics/assumptions:

The model's lowest layer is designated as PBL

The PBL depth is predicted

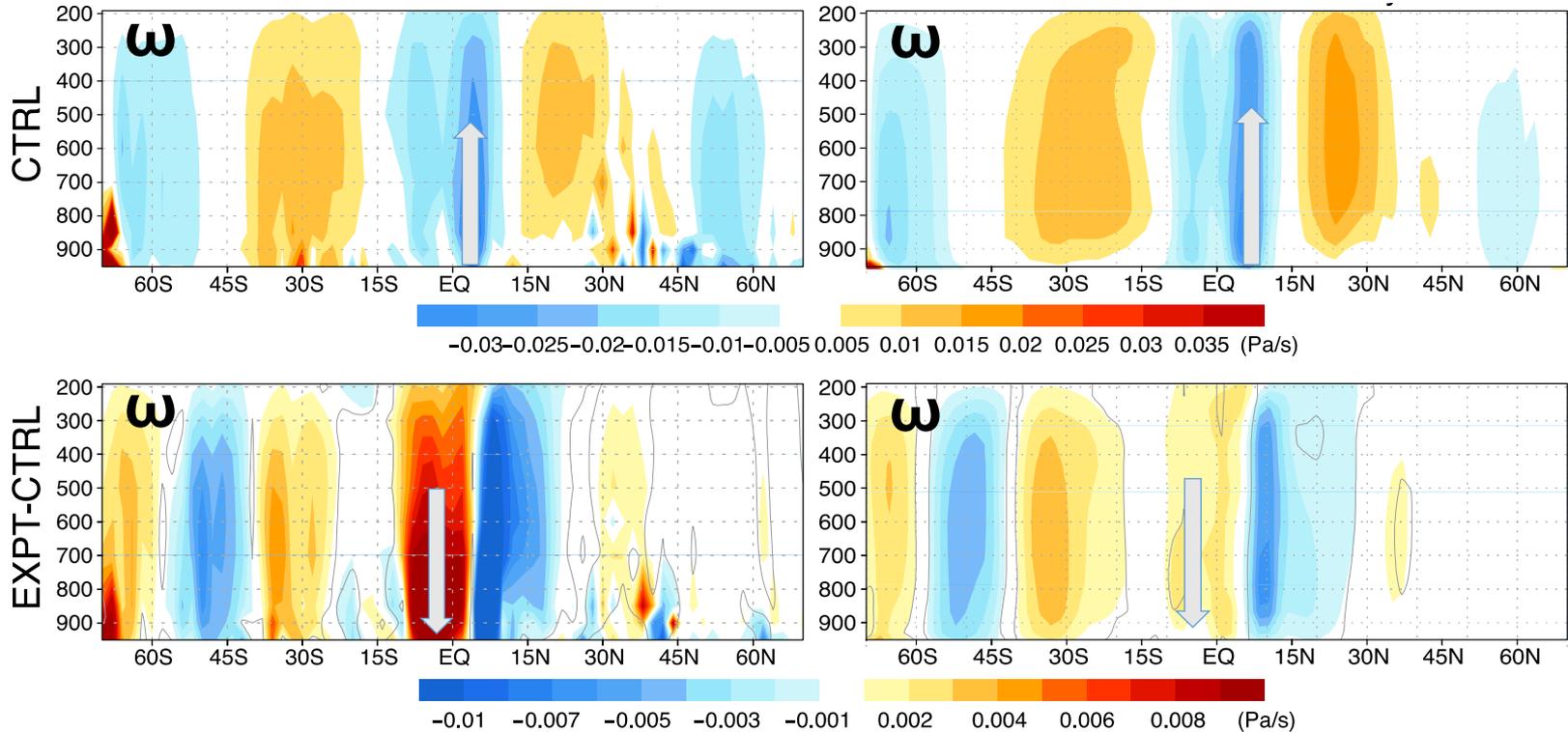
Stratocumulus (PBL-top clouds) are determined implicitly

The bottom of cumulus clouds is at the PBL-top

In EXPT subsidence increases in southern tropics, more so in the UCLA CGCM

CGCM

ESM



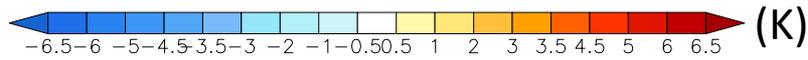
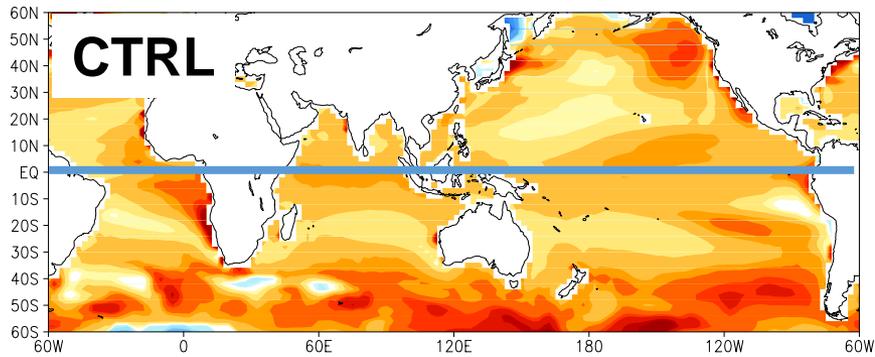
Top: Control

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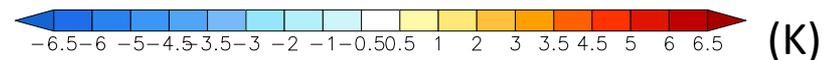
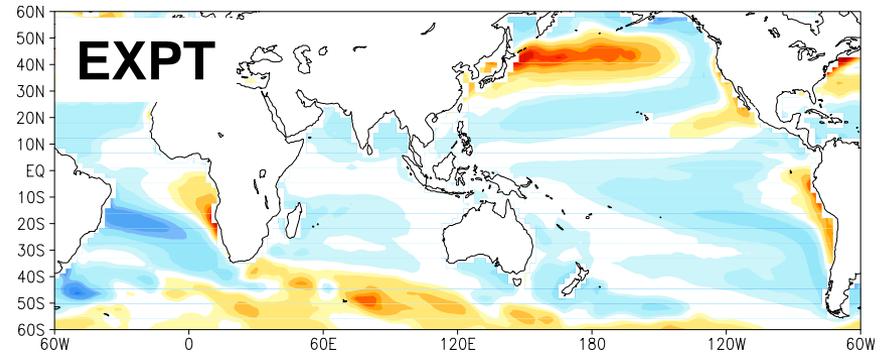
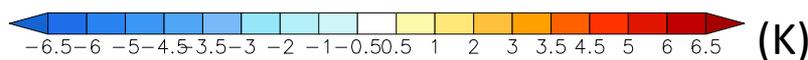
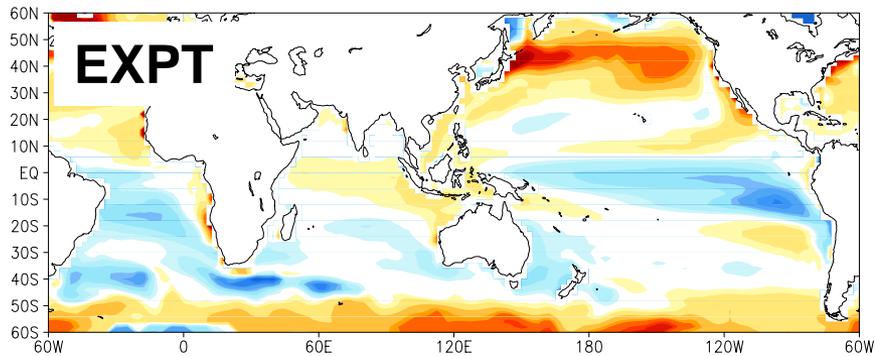
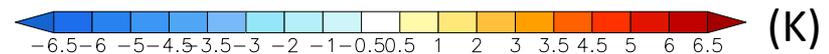
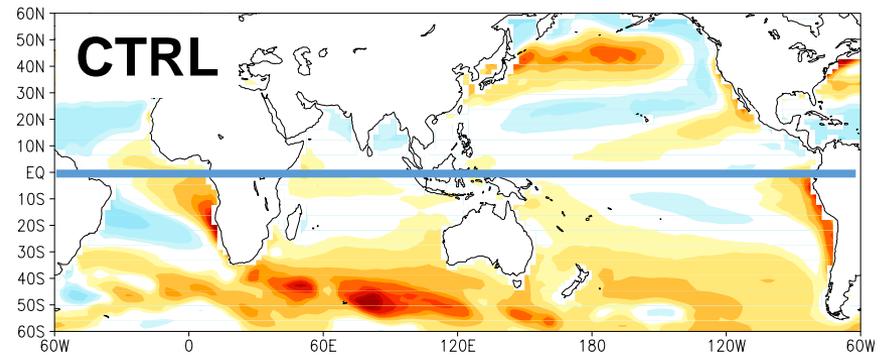
Also, the ITCZ is enhanced in northern tropics

SST biases in EXPT are generally weaker in both models!

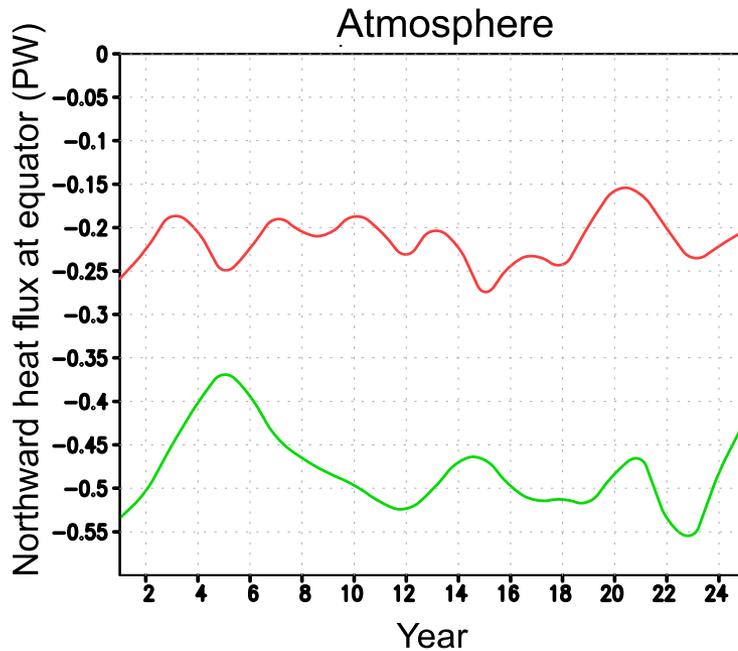
CGCM



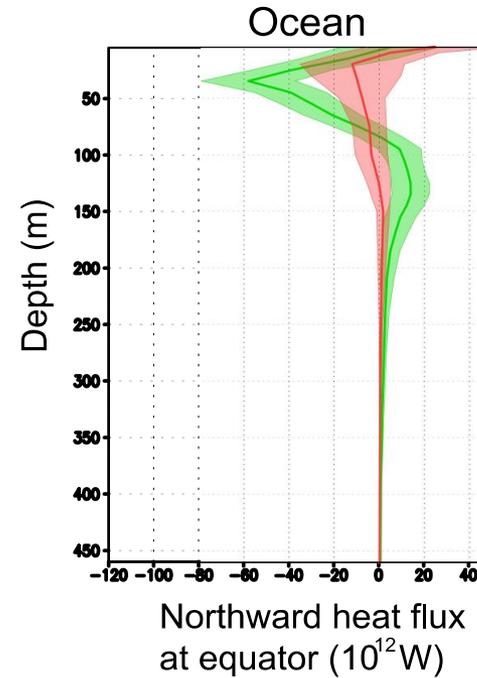
ESM



Change in northward heat flux at the equator in the atmosphere and ocean in the experiment.



CGCM



ESM